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Motivation

Although the subarctic tundra represents a large fraction of the Earth's surface, very little climatological data are available for the validation of models used to anticipate how these changes will affect weather patterns and local hydrological processes.

This project aims at obtaining a comprehensive climatological and hydrological image of an archetype of such an environment near Umiujaq in northern Québec, Canada. To achieve this, fluxes of water vapor and heat between soil, snow and atmosphere are traced and subsequently energy and water budgets are calculated from these fluxes.

Here, data from summer and fall 2017 is presented.



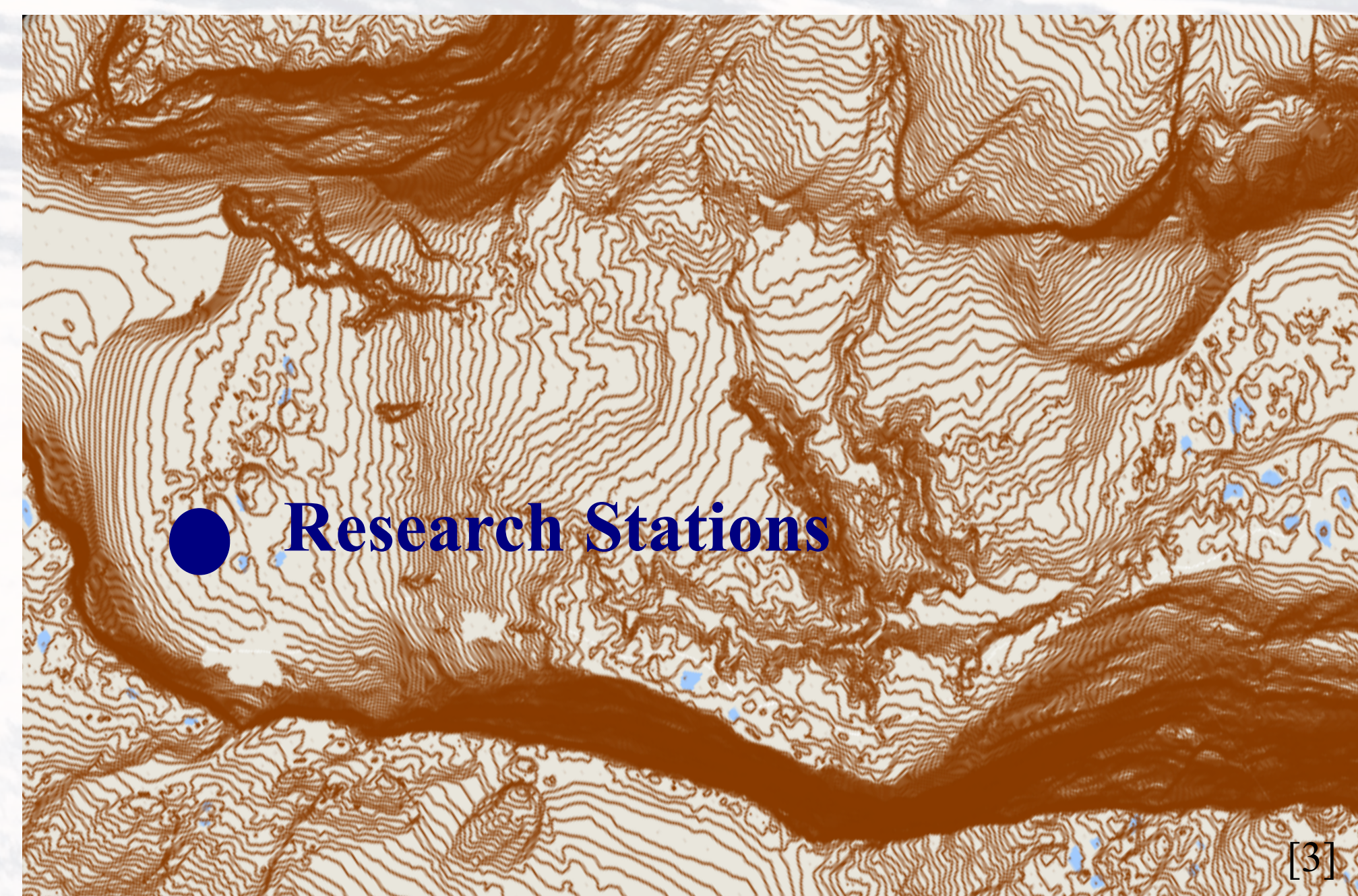
The Tasiapik Valley

Site Description

The study site is a 2.1 km² watershed near Umiujaq, which exhibits discontinuous permafrost and a medium slope confined by cuestas. The vegetation mainly consists of shrubs but patches of lichen and forest can be found as well.

Climatology

Mean annual temp.: -3°C; Mean temp. February: -22.4°C; Mean temp. June: 12.2°C
Total precipitation: 500-600 mm
Snowmelt: May - June; Mean wind speed: 5.9 m/s
Max. wind speed 28 m/s; Dominant wind dir. - South.

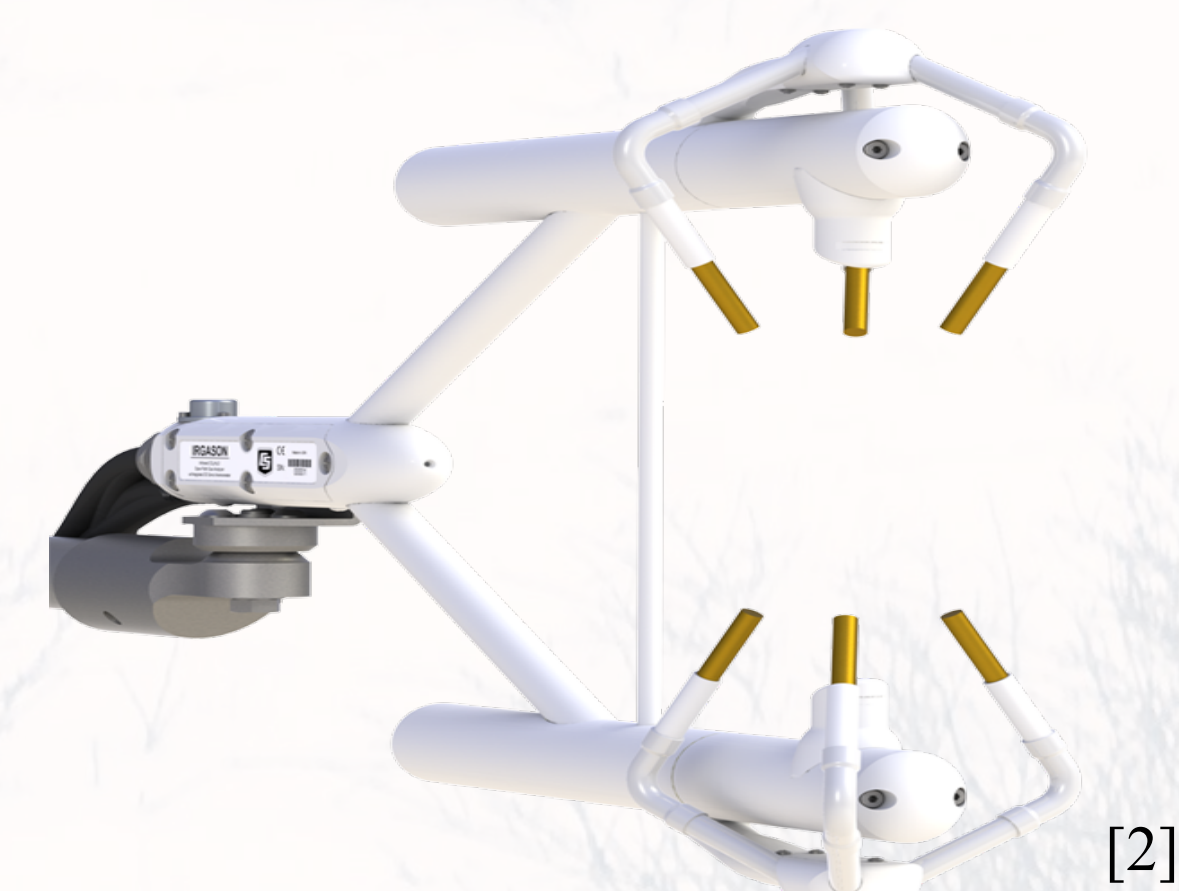


Eddy Covariance

Turbulent flux calculation with 3D wind speed and H₂O/CO₂ concentration measurements

- + High frequency
- + Non disturbing
- + Continuous
- Requires turbulence
- Gap filling

Eddy Covariance System



The deployed eddy covariance system, the IRGASON [2]

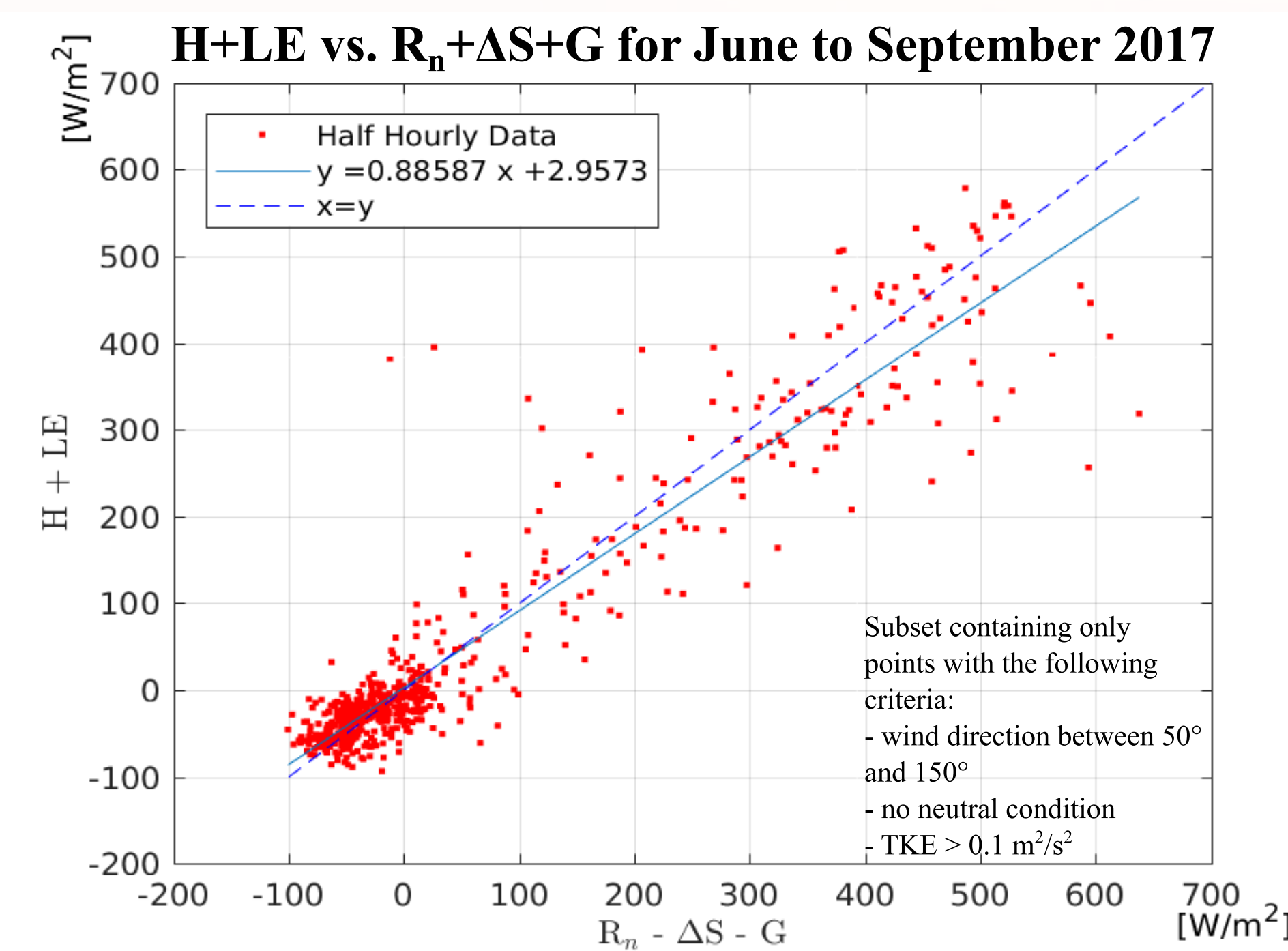
Research Stations



Left: Geonor precipitation measurement
Middle: SILVA tower with eddy covariance measurements
Right: Tundra station with radiation measurements

Methodology

Energy Budget



$$R_n = H + LE + G + \Delta S$$

R_n: Net Radiation from a CNR4 radiometer

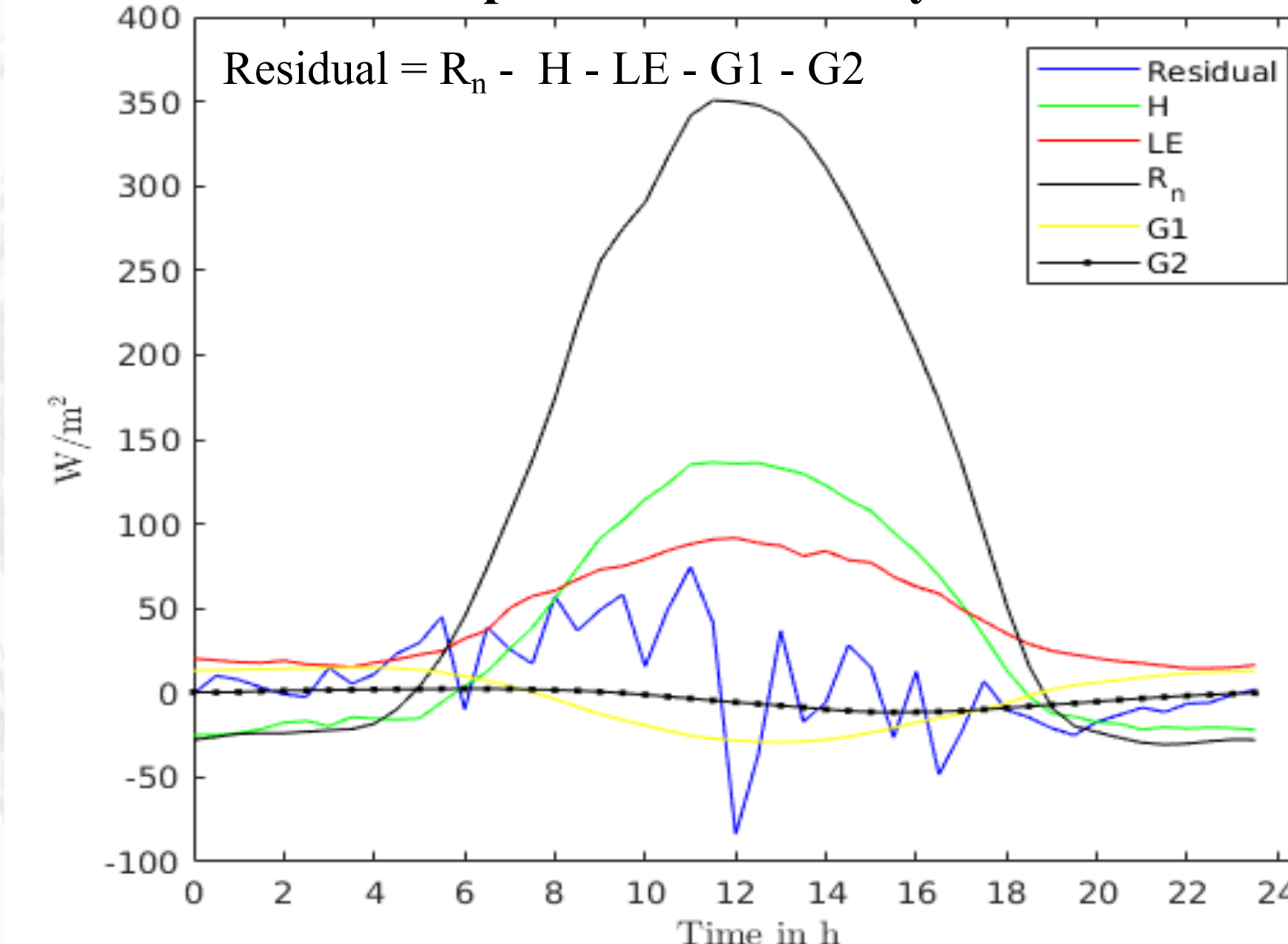
H: Sensible Heat from the IRGASON

LE: Latent Heat from the IRGASON

G: Ground Heat Flux derived from thermocouples (G1) and ground heat flux plates (G2)

ΔS: Storage Terms of the vegetation

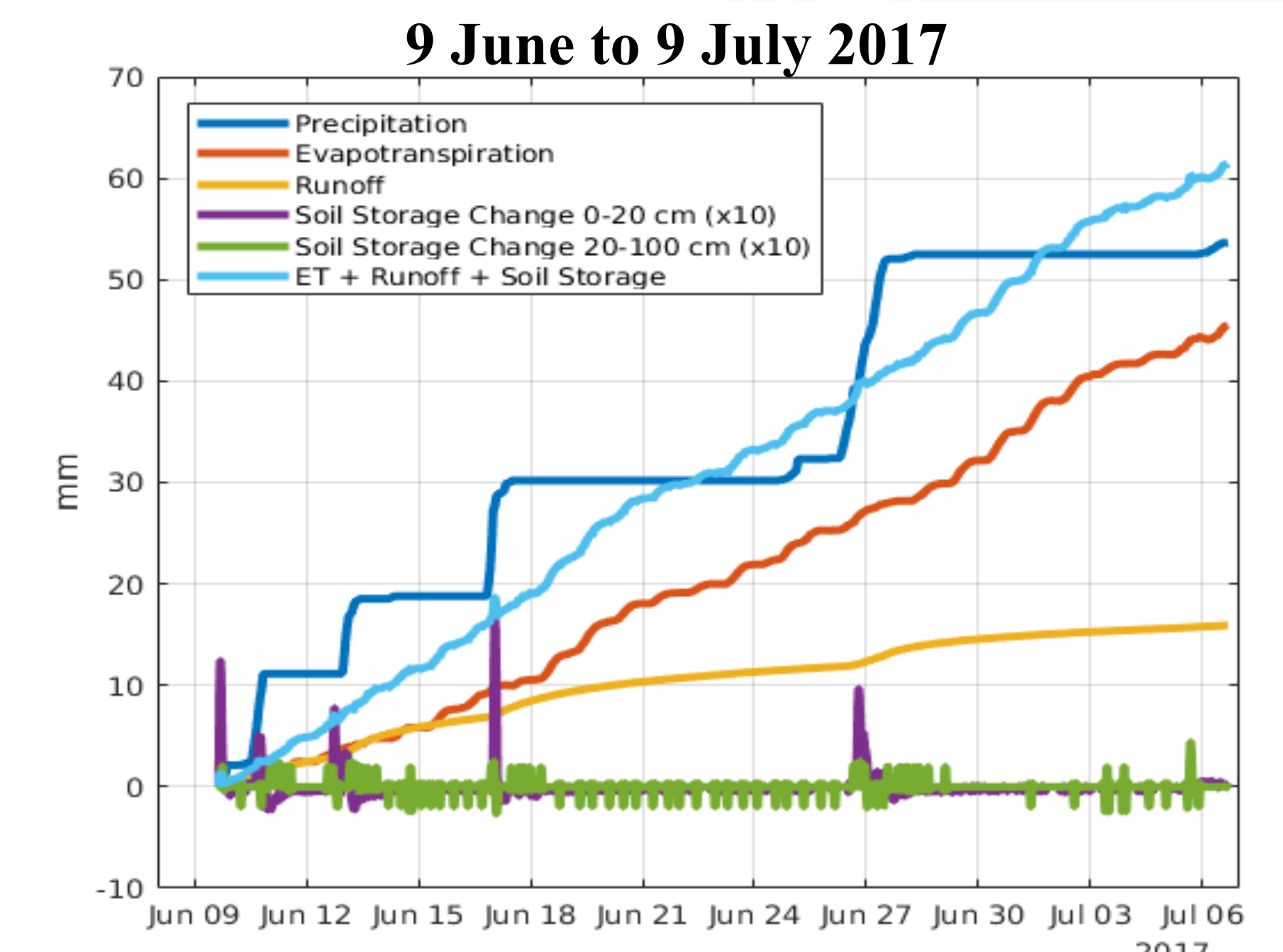
Mean pattern of the daily residual



Conclusions

- Underestimation of H and LE during daytime
- Energy closure of: 88.5%
- ➔ More energy available than being consumed by H and LE
- Main reasons for the underestimation of H and LE:
 - ➔ Complex terrain
 - ➔ Advection
 - ➔ Uncertainties in the ground storage change
 - ➔ Missing vegetation heat storage

Water Budget



$$P = R + ET + \Delta S$$

P: Precipitation measured by a rain gauge

R: River runoff through a flume

ET: Evapotranspiration from the IRGASON

S: Water storage change in soil, lakes and groundwater

Conclusions

- ET is ≈79% higher than runoff on average
- ➔ ET is the driving force of water loss during summer
- Minor contribution of the upper soil layer to the water storage
- ➔ Groundwater flows are the dominant storage terms
- ET + R is higher than the precipitation
- ➔ Dry out during this one month period

Future Work

- Examine the effect of the coordinate system on the energy balance closure
- Estimate storage terms of the vegetation and incorporate them into the energy balance
- Include terms of groundwater flow and water storage in lakes into the water budget
- Extend work to winter and include snow processes to analysis

Acknowledgements

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Sources

[1] © User:Carport / Wikimedia Commons / CC-BY-SA-3.0

[2] Cambell Scientific [3] Geoindex, Université Laval

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